



# *Cylindroporella sugdeni* Elliott, 1957, an Early Cretaceous Middle Eastern Dasycladalean alga – a revision

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**Abstract.** The type material of *Cylindroporella sugdeni* Elliott, 1957, is revised. The total lack of sterile laterals and consequently the absence of the diagnostic feature of the genus *Cylindroporella*, i.e., the alternation of sterile and fertile laterals within the same whorl, leads to its exclusion from the genus *Cylindroporella* Johnson, 1954, and to its ascription to the genus *Holosporella* Pia, 1930. Recent stratigraphic investigations suggest that the stratum typicum originally reported as Early Cretaceous in age is more precisely Hauterivian or possibly Early Barremian in age.

## 1 Introduction

Except for the fossil species ascribed to the fossil family Setonellaceae in which the lateral expansions are randomly distributed along the alga main axis, the Dasycladalean algae have laterals arranged in regular rows that are called “verticils”. The latter comprise either sterile laterals or fertile laterals, except for the representatives of *Cylindroporella* Johnson, 1954, and *Korkyrella* Sokač & Velić ex Sokač, 2004, both ascribed to the fossil family Triploporellaceae. Having their sterile and fertile laterals occurring in alternation within the same row, both genera have a unique body plan amidst the fossil and living Dasycladalean algae. The more their representatives are reassessed, the more one questions the existence of such a body plan (see Barattolo and Parente, 2000, for instance).

Following the description of its type species, *Cylindroporella barnesii* Johnson, 1954, two new species were introduced by Elliott (1957): the Early Cretaceous *C. sugdeni* and the Late Jurassic *C. arabica*. The revision of the second species listed, i.e., *C. sugdeni* Elliott, 1957, is addressed below.

## 2 Material and methods

Thanks to a loan organized by the Department of Palaeobotany of the Natural History Museum, London, the five thin sections that correspond to the type material of *Cylindroporella sugdeni* Elliott, 1957, which are currently stored in the Collection PAL (Palaeontology) under numbers PB V 41620 to 41624 (Fig. 1), were examined. The thin section labels indicate that the material comes from a depth of 4380', i.e., 1335 m (measured depth, MD), from core 14 in well Fahud 1 in Oman.

In addition to this material, another thin section, which is part of a set given by Graham F. Elliott to Jacques Emberger (Université de Bordeaux), was also studied. The label of this thin section implies it comes from the well “Fahud 1 (Oman), core 14, depth 4381' (MD)”. This last thin section is currently deposited with a LPB (Laboratoire de Paléontologie de Brest) number in the collections of the Département des Sciences de la Terre et de l'Univers, Université de Bretagne Occidentale, Brest (France).

The dominant texture is a bioclastic grainstone with micriticized grains of all sizes (peloids to aggregates, micritic extraclasts) and a few ooids. The skeletal grains consist mostly of the new alga, indeterminate lituolid and textulariid



**Figure 1.** The set of six thin sections studied. Both sides of 41623 and 41624 are visible; the back sides of them were rotated in order to make the text readable.

foraminifers, rare *Praechrysalidina infracretacea* Luperto Sinni, 1979, and echinoderm remains. Unfortunately, the thin sections are rather thick (probably thicker than 30 µm). The original piece of limestone was porous but it was not injected with resin in order to consolidate it. As a consequence the material is locally crumbled (with local overlaps) or was partly removed. In addition, small air bubbles are trapped in the Canada balsam. In order to better understand the alga structure, a photo editor was used to color in blue the holes in the thin section views (Fig. 2); then some visible features (Fig. 3), possibly belonging to the alga, were contoured.

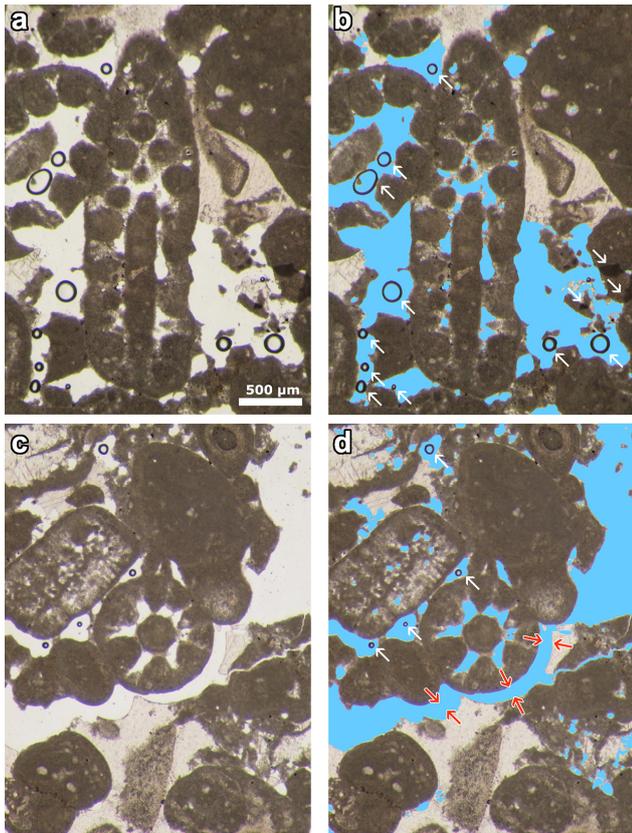
### 3 Geological background

#### 3.1 Type locality

Well Fahud 1 (coordinates: 22°17'25.4" N, 56°30'52.6" E) was “spudded-in” on 18 January 1956. According to Morton (2006, p. 191), “Drilling was completed at a depth of 3760 m in the Paleozoic evaporites, probably Cambrian, on 28 May 1957”. From ca. 850 to ca. 1650 m depth, it drilled through the Lower Cretaceous Kahmah Group of Oman (Fig. 4).

#### 3.2 Stratigraphical background

The name of the Lower Cretaceous group changes across the international boundary. In the United Arab Emirates (UAE,



**Figure 2.** (a–b) Oblique section, paratype (one of the syntypes), thin section PB V 41624 (Elliott, 1957, pl. 1, fig. 6); (c–d) transverse section, paratype (one of the syntypes), thin section PB V 41623 (Elliott, 1957, pl. 1, fig. 3). The parts colored in blue correspond to parts removed while making the thin section and possibly some original porosity. White arrows pointing to the upper right indicate air bubbles and those pointing to the lower left indicate overlaps in the thin section manufacturing; opposing red arrows indicate an induced fracture. All photos on the same scale (scale bar = 500 µm).

among which is the emirate of Abu Dhabi) and in Saudi Arabia, the equivalent Lower Cretaceous unit is called the Thamama Group (Fig. 4). Core 14 with *Cylindroporella sugdeni* Elliott, 1957, was retrieved from an interval labeled as “Habshan”. For decades, oil and gas company geologists have developed and used an interregional stratigraphic framework for this interval (e.g., Hassan et al., 1975). According to Granier (2000; Granier et al., 2003), most of their operational units prove to be unconformity-bounded units. In addition, because they can also be characterized by their discrete fossil contents, Granier (2008) treats them as genuine regional stages. In Abu Dhabi, the shallow-water carbonate units of the Thamama Group comprise, from top to bottom, the Bab and the Shu’aiba (Upper Aptian), the Hawar (Lower Aptian), the Kharaiib (Upper Barremian–Lower Aptian), the Lekhwair (Upper Valanginian–Hauterivian–Lower Barremian), the Zakum (Lower Valanginian), the Belbazem

(Berriasian), the Bu Haseer (transition of the Tithonian to the Berriasian), and the Habshan (Tithonian). However, in Oman, the shallow-water carbonates of the lower part of the interval may be missing or they may pass to hemipelagic and pelagic facies of the Kahmah Group (Granier, 2000, 2008; Granier et al., 2003). In addition to the UAE units, two facies-driven formations were originally described from the Kahmah Group in Oman: the Rayda and the Salil (e.g., see Hassan et al., 1975). The Rayda, which is also an unconformity-bounded unit, is time equivalent to both the Bu Haseer and Belbazem (Tithonian–Berriasian). The Salil facies is the basinal time equivalent to the rest of the stratigraphic column (Valanginian–Aptian pars). Accordingly, there are not only geopolitical reasons but also geological reasons (ages, hiatuses, and development of basinal facies) that may justify the use of the Thamama Group in the UAE and that of the Kahmah Group in Oman.

### 3.3 Stratigraphic location of the type level

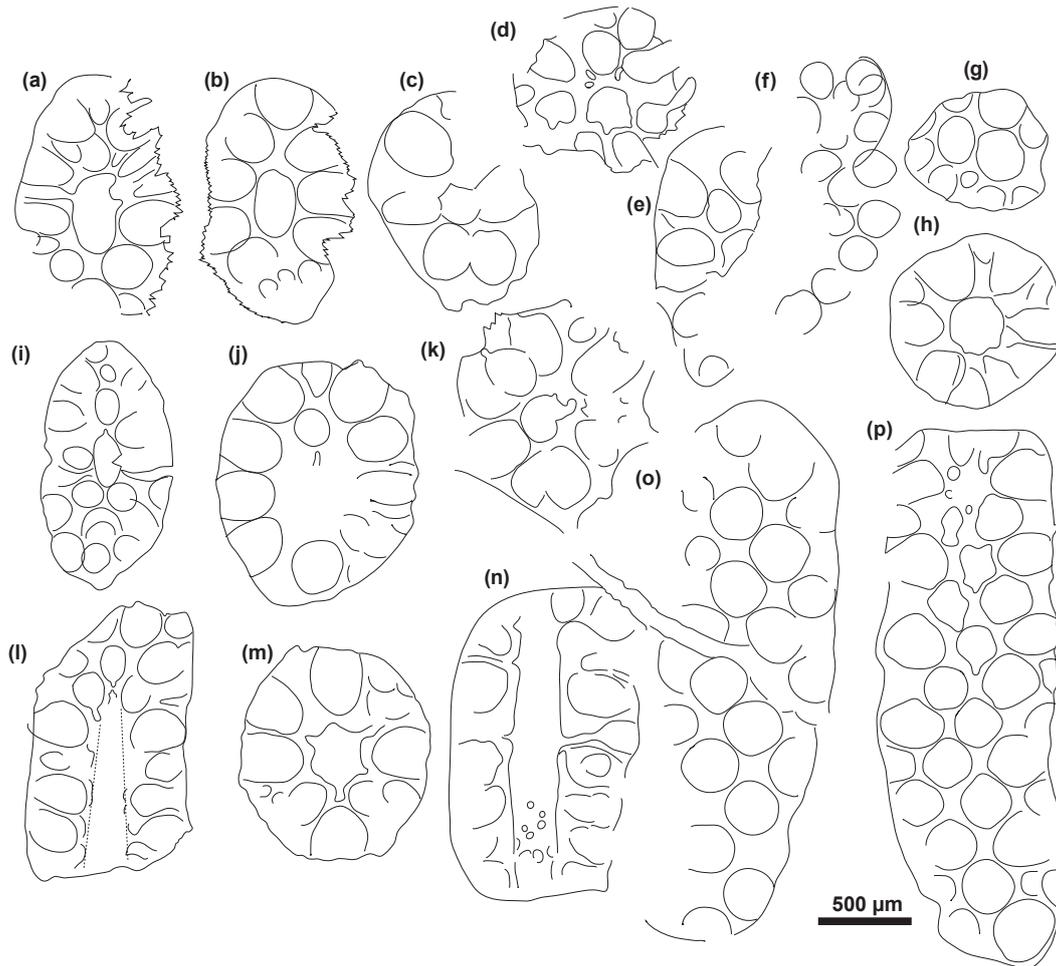
In well Fahud 1, the lithostratigraphic succession identified by Petroleum Development Oman (PDO) and Shell geologists comprises, from top to bottom, the “Shu’aiba” (that includes the Hawar), the Kharaiib, the “Lekhwair”, the “Habshan”, and the Salil (Fig. 4). The use of quotation marks is justified by the fact that the definition of these units does not match the regional standardized scheme (Hassan et al., 1975; Granier 2000, 2008; Granier et al., 2003). This is especially true in the case of the Habshan of the well Fahud 1, which refers to a transitional zone between the Salil basinal facies below and Lekhwair facies above and which has no connection with the type Habshan in Abu Dhabi (latest Jurassic in age). In conclusion, because the type level of *Cylindroporella sugdeni* Elliott, 1957, is located in the median part of an interval equivalent to the type Lekhwair (it is not located in the Kharaiib as erroneously stated in Granier, 2008), and although an Early Barremian age cannot be excluded, it is most probably Hauterivian in age (Fig. 4).

## 4 Systematic palaeontology

### 4.1 Comparison with other species

As stated in the introduction above, there are only two genera the sterile and fertile laterals of which occur in alternation within the same row: *Korkyrella* Sokač & Velić ex Sokač, 2004 and *Cylindroporella* Johnson, 1954.

*Korkyrella* Sokač & Velić ex Sokač, 2004, which is monospecific, is based on *K. texana* Johnson ex Sokač, 2004 (“*Salpingoporella texana* Johnson, 1965”, nom. nud.). On the one hand, *Cylindroporella* is also a valid name for a Bryozoan genus, *Cylindroporella* Hincks, 1877, based on *Lepralia tubulosa* Norman, 1868, which is its sole representative. On the other hand, some years ago there were still nearly 20 species referring to the algal genus *Cylindroporella* Johnson,



**Figure 3.** Contours of some algal remains. (a–b, g–h, j, l) Thin section PB V 41621; (c, f, i, m) thin section PB V 41622; (d–e, k, n–p) thin section PB V 41620. (d) Syntype (Elliott, 1957, pl. 1, fig. 1); (h) syntype (Elliott, 1957, pl. 1, fig. 3); (l) syntype (Elliott, 1957, pl. 1, fig. 2); (m) syntype (Elliott, 1957, pl. 1, fig. 4). All drawings on the same scale (scale bar = 500 μm).

1954, which is based on *Cylindroporella barnesii* Johnson, 1954. Today, due to some recent revisions, the number of its representatives has significantly diminished (Table 1).

According to Sokač (2004), sterile laterals of *Korkyrella texana* Johnson ex Sokač, 2004, are rather long and thin primaries; the fertile laterals consist of ovoid primaries followed by short secondaries. According to Conrad (1982), sterile laterals of *Cylindroporella barnesii* Johnson, 1954 consist of rather long and thin primaries followed by short secondaries; the fertile laterals are ovoid primaries. However, *Cylindroporella sugdeni* Elliott, 1957 does not match any of these diagnostic criteria. There are not any long and thin laterals, supposedly sterile, but only ovoid primaries without secondaries, supposedly fertile. In tangential sections, the empty spaces remaining between adjacent ovoid primaries could have been confused with sterile primaries. In this case, the algal sections may recall those of *Montiella elitzae* (Bakalova, 1971), a form with laterals consisting of a thin primary bearing a thin secondary and an ovoid am-

pulla at its outer end. Nonetheless, several photomicrographs clearly attest the lack of fertiles. Consequently, the species is transferred to the genus *Holosporella* (Pia, 1930), which is characterized by a simple body plan with a main axis bearing closely set verticils of vesicular laterals, commonly arranged in a quincunx from one verticil to the next.

### Phylum **Chlorophyta**

Class **Dasycladophyceae** Hoek et al.

Order **Dasycladales** Pascher

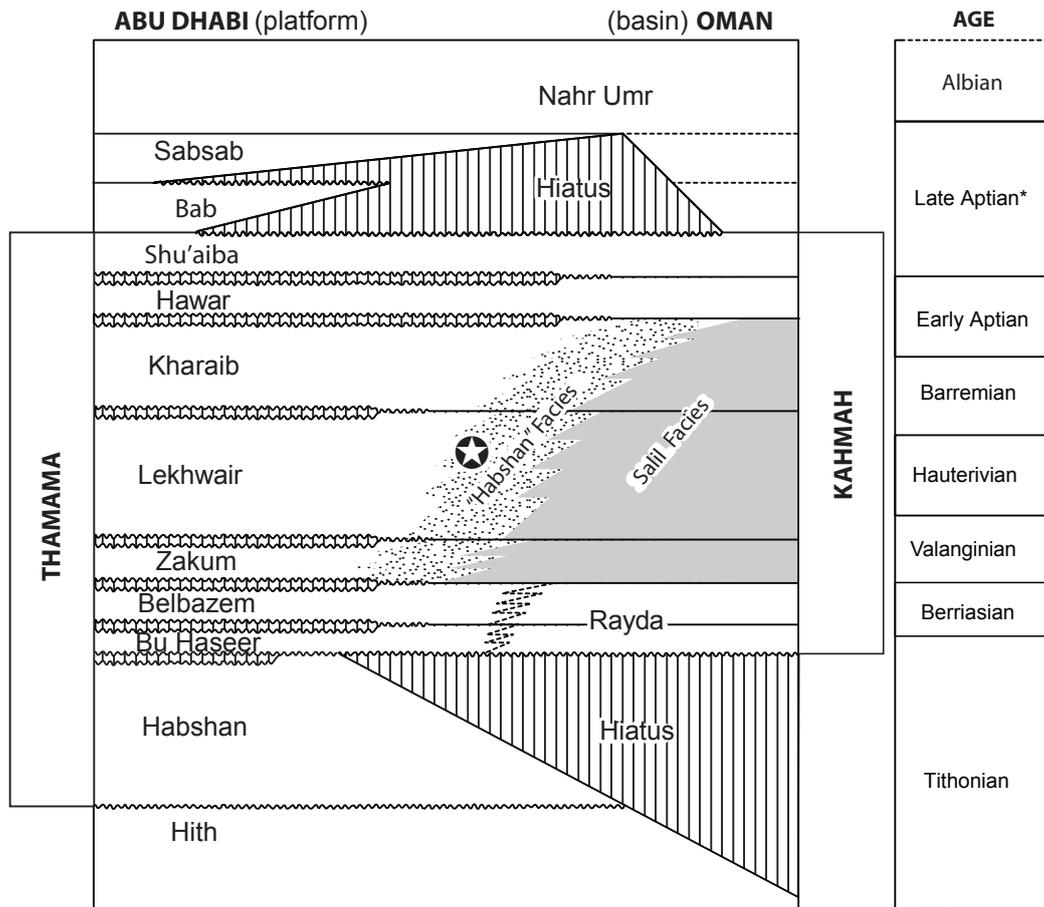
Family **Triploporellaceae** Pia, 1920

Tribe **Salpingoporelleae** Bassoullet et al., 1979

Genus *Holosporella* Pia, 1930

*Holosporella sugdeni* Elliott, 1957, nov. comb.

(Figs. 2–3 and 5)



**Figure 4.** Stratigraphic framework of the United Arab Emirates and Oman (modified from Granier et al., 2003; Granier, 2008; Granier and Busnardo, 2013; \* here the Late Aptian includes the Furcata ammonite Zone). The large star indicates the supposed relative location of the core sampling.

1957 *Cylindroporella sugdeni* Elliott: 227, pl. 1, figs. 1–6.

1968 *Cylindroporella barnesii* Johnson; Elliott: 37, pl. 6, fig. 3.

1968 *Cylindroporella sugdeni* Elliott; Elliott: 38, pl. 6, figs. 5 (Elliott, 1957, pl. 1, fig. 2), 6 (Elliott 1957, pl. 1, fig. 5), and 7 (Elliott, 1957, pl. 1, fig. 1).

1968 *Cylindroporella sugdeni* Elliott; Johnson: 36, pl. 6, fig. 4.

1969 *Cylindroporella sugdeni* Elliott; Johnson: 56, 153, pl. 41, figs. 1 (Elliott, 1957, pl. 1, fig. 1), 2 (Elliott, 1957, pl. 1, fig. 5), 3 (Elliott, 1957, pl. 1, fig. 2), 4 (Elliott, 1957, pl. 1, fig. 4), and 5 (Elliott, 1957, pl. 1, fig. 6).

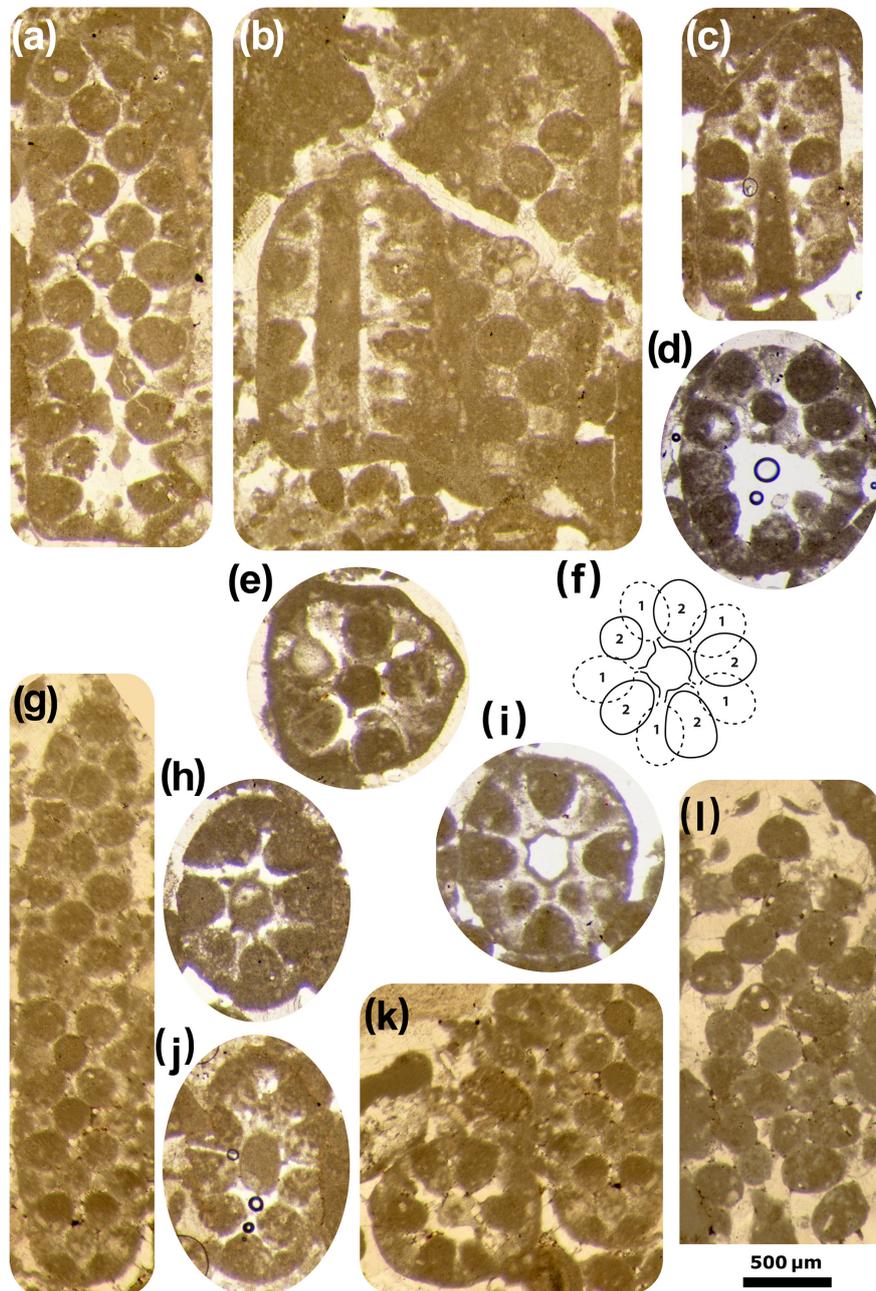
non 1971 *Cylindroporella sugdeni* Elliott; Basson & Edgell: 417, pl. 2, figs. 1–3 (*Montiella elitzae*; Bakalova, 1971).

1971 *Cylindroporella sugdeni* Elliott; Basson & Edgell: 417, pl. 2, fig. 4.

1978 *Cylindroporella sugdeni* Elliott; Bassoulet et al.: 179–181, pl. 8, figs. 1 (Elliott, 1957, pl. 1, fig. 1), 2 (Elliott, 1957, pl. 1, fig. 2), and 3.

**Original diagnosis:** (Elliott, 1957) “Large, thick cylindroporelliform segments, agreeing in internal structure with those of *Cylindroporella barnesii*, but with greater diameter of central canal and sporangia and externally, and a greater distance between whorls; it is proportionally shorter and much thicker than the type species (...).”

**Description:** Thalli are found in the form of broken calcareous tubes with an axial cavity in lieu of the main axis and globular pores in lieu of the laterals. The inner diameter of such a calcareous coating commonly represents one-fourth of its outer diameter. The calcification may vary from one specimen to the others or within a single specimen. Usually it forms a thin crust around the laterals and the main axis. Micritic infills may be observed above and beyond this crust (Fig. 5b right, g and k right) but in some cases the “intercellular” space, i.e., the empty space remaining



**Figure 5.** (a) Tangential section, paratype (one of the syntypes), thin section PB V 41620 (Fig. 3p; Elliott, 1957, pl. 1, fig. 1); (b) axial (left) and tangential (right) sections, thin section PB V 41620 (Fig. 3n–o); (c) oblique section, lectotype (one of the syntypes), thin section PB V 41621 (Fig. 3l; Elliott, 1957, pl. 1, fig. 2); (d) oblique section, thin section PB V 41621 (Fig. 3j); (e) transverse section, paratype (one of the syntypes), thin section PB V 41623 (Elliott, 1957, pl. 1, fig. 5); (f) sketch of photomicrograph 5e illustrating the imbrication of two whorls. The solid lines correspond to the laterals of one whorl (2) and the main axis with the peduncular parts of the laterals of the next whorl. The dotted lines correspond to the laterals of the next whorl (1); (g) tangential section, thin section 4381, Collection Emberger, Brest; (h) sub-transverse, slightly oblique section, thin section PB V 41624; (i) sub-transverse, slightly oblique section, paratype (one of the syntypes), thin section PB V 41622 (Fig. 3m; Elliott, 1957, pl. 1, fig. 4); (j) slightly oblique section, thin section PB V 41621; (k) axial (left) and tangential (right) sections, thin section 4381, Collection Emberger, Brest; (l) subtangential, slightly oblique section (the main axis is visible at the bottom of the picture), Collection Emberger, Brest. All photos on the same scale (scale bar = 500 µm).

**Table 1.** Nomenclatural changes that have affected more than half of the former representatives of the genus *Cylindroporella* Johnson, 1954.

Species referred to the genus <i>Cylindroporella</i> Johnson, 1954	New status (genus) after revision or new combination	New status (family) after revision or new combination	Author(s) of the revision or new combination	Year of the revision or new combination
<i>C. elitzae</i> Bakalova, 1971	<i>Montiella</i> (Munier-Chalmas ex L. & J. Morellet, 1922)	Dasycladaceae	Radoičić	1980
<i>C. benizarensis</i> Fourcade et al. ex Jaffrezo in Bassoullet et al., 1978, non 1972	<i>Montiella</i> (Munier-Chalmas ex L. & J. Morellet, 1922)	Dasycladaceae	Granier	1990
<i>C. barbui</i> Dragastan, 1989, non 1978	Junior synonym of <i>Montiella elitzae</i> (Bakalova, 1971)	Dasycladaceae	Conrad & Varol	1990
<i>C. adducta</i> Maslov, 1960	Possible senior synonym of <i>Montiella elitzae</i> (Bakalova, 1971)	Dasycladaceae	Bucur	1993
“ <i>C. anici</i> Nikler & Sokač, 1965” (nom. nud. fide Granier & Deloffre, 1993)	<i>Heteroporella</i> ? <i>anici</i> Nikler & Sokač in Granier & Deloffre, 1993, non-1965	left in open nomenclature	Granier et al.	1994
<i>C. ellenbergeri</i> Lebouché & Lemoine in Granier & Deloffre, 1993, non-1963	<i>Chinianella</i> Ott ex Granier et al., 1994, non-1967	Polyphysaceae	Granier et al.	1994
<i>C. cruciformis</i> Granier & Brun, 1991	<i>Fourcadella</i> n. gen. Granier, 2002	Thyrsoporellaceae	Granier	2002
<i>C. parva</i> Radoičić, 1983	<i>Montiella</i> (Munier-Chalmas ex L. & J. Morellet, 1922)	Dasycladaceae	Radoičić	2006
“ <i>C. maslovi</i> Srivastava, 1973” (nom. nud. fide Granier & Deloffre, 1993)	Junior synonym of <i>Montiella elitzae</i> (Bakalova, 1971)	Dasycladaceae	Bucur	2011
<i>C. elassonos</i> Johnson & Kaska, 1965	It is not an alga, but a foraminifer		Granier et al.	2013
<i>C. lusitanica</i> Ramalho, 1970	<i>Barattoloporellopsis</i> n. gen. Granier et al., 2017	Dasycladaceae	Granier et al.	2017
<i>C. sugdeni</i> Elliott, 1957	<i>Holosporella</i> Pia, 1930	Triploporellaceae	this work	

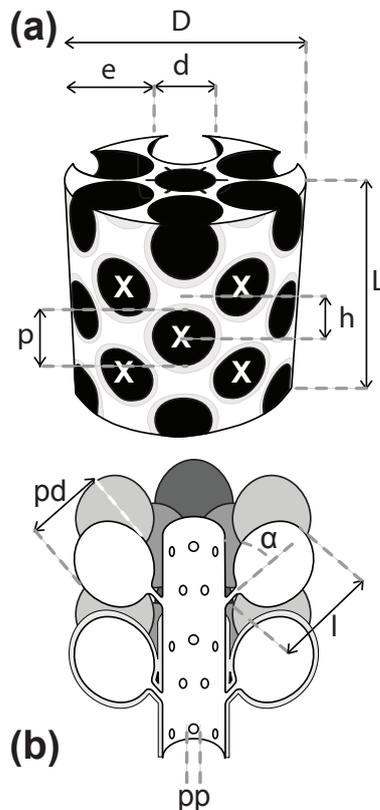
between the vesicular laterals and the main axis, is filled in by a drusy calcite cement (Fig. 5a and l). Previously, such intercellular micritic infills with ill-defined contours were erroneously interpreted as thin sterile laterals.

No unquestionably cyst-like structures were observed but the laterals are supposedly fertile because their volume is significantly larger than that of the axis cavity, which excludes that the species is endosporous (with cysts located in the main axis) and supports it being interpreted as cladospore (with cysts located in the primary laterals). The proximal part of these laterals consists of a short and narrow peduncle (Fig. 5c, h–j and k left) that expands abruptly into the characteristic wider vesicular part.

Ovoid laterals from successive verticils have a quincunx arrangement (Fig. 6a) that results from the imbrication of the verticils, clearly visible in tangential sections (Fig. 5a,

b right, g, k right and l). The imbrication ratio is computed by the difference between the maximum width of a lateral, i.e., the thickness of a verticil, and the distance between two successive verticils at their insertion points, divided by this last distance. Imbrication is particularly strong in this species in which it represents nearly two-thirds of the whorl spacing.

**Emended diagnosis:** Rather large *Holosporella* species with a relatively narrow main axis and strongly imbricated verticils of wide vesicular laterals, supposedly fertile. Using the various measurements listed below ( $D$ ,  $d$ ,  $d/D$ ,  $h$ , etc.) may help discriminate it from the other representatives of the genus.



**Figure 6.** Reconstructions (a) of a heavily calcified *Holosporella* and (b) of a weakly calcified one. In (a) the five X's at the center of the pores illustrate the quincunx arrangement of the laterals. Key for the measurements is as follows. *L*: maximum length observed; *D*: outer diameter; *d*: inner diameter; *h*: whorl spacing; *e*: thickness of the calcareous coating;  $\alpha$ : angle of the pore to the main axis; *l*: length of the primary pores; *p*: visible width of a primary pore; *pp*: width of the proximal part of the primary pores; *pd*: maximum width of the primary pores.

**Lectotype:** The lectotype (Figs. 3l and 5c) selected here is the oblique section of pl. 1, fig. 2 of Elliott (1957) from the thin section labeled PB V 41621.

**Remark:** Some features observed in axial sections are not visible in transverse sections and vice versa. Oblique sections should always be preferred as a type (holotype, lectotype, neotype) to axial and transverse sections because a palaeophycologist usually gathers more information from oblique sections than from any axial or transverse section.

**Paratypes:** The paratypes are the five remaining specimens: pl. 1, fig. 1 of Elliott (1957) from the thin section labeled PB V 41620 (Fig. 5a); pl. 1, fig. 3 of Elliott (1957) from the thin section labeled PB V 41621 (Figs. 2c–d and 3h); pl. 1, fig. 4 of Elliott (1957) from the thin section labeled PB V 41622 (Figs. 3m and 5i); pl. 1, fig. 5 of Elliott (1957) from the thin section labeled PB V 41623 (Fig. 5e–f); and

pl. 1, fig. 6 of Elliott (1957) from the thin section labeled PB V 41624 (Fig. 2a–b).

#### Measurements: (Fig. 6)

- *L* (maximum length observed) = 5.8 mm versus > 3.0 mm for Elliott (1957);
- *D* (outer diameter) =  $1.07 \pm 0.12$  mm (18 measurements), minimum = 0.84 mm, maximum = 1.26 mm versus 0.78–1.14 mm for Elliott (1957);
- *d* (inner diameter) =  $0.27 \pm 0.05$  mm (18), minimum = 0.18 mm, maximum = 0.35 mm versus 0.23–0.36 mm for Elliott (1957);
- $d/D = 25.0 \pm 3.7\%$  (18)
- *h* (whorl spacing) = 0.20 mm in average (seven measurements) versus 0.39 mm for Elliott (1957) who erroneously did not measure the distance of two successive whorls, but instead the distance of two whorls with an intermediate one;
- *e* (thickness of the calcareous coating) =  $D - 2d = 0.53 \pm 0.10$  mm (18);
- $\alpha$  (angle of the pore to the main axis) = ca.  $60^\circ$ ;
- *l* (length of the primary pores) =  $e / \sin \alpha =$  ca. 0.56 mm;
- *pp* (width of the proximal part of the primary pores) = 0.04 mm;
- *pd* (maximum width of the primary pores) = 0.33 mm versus 0.31 mm for Elliott (1957);
- *i* (imbrication of the whorls) =  $pd - h = 0.13$  mm, a positive value;
- $i/h$  (imbrication ratio) =  $(pd - h)/h =$  ca. 65 %;
- *w* (number of pores per verticil) = 5 to 8.

## 5 Conclusions

*Holosporella sugdeni* Elliott, 1957, nov. comb., is known with certainty at least from the Lekhwaier and Kharaib of Abu Dhabi (Granier et al., 2003) and from the Jezzian of Lebanon (Basson and Edgell, 1971), i.e., from the Hauterivian and Barremian interval in the Middle East. More than half a century after it was introduced by Elliott (1957), there are few correct identifications of this species. Records outside this region commonly come with “cf.” or “aff.” and do not expressly refer to this species but most probably to *Montiella elitzae* (Bakalova, 1971). Thus, *Holosporella sugdeni* Elliott, 1957, nov. comb., might be an endemic species.

Following the revision of this species, the genus *Cylindroporella* Johnson, 1954, has lost another representative. As

Barattolo and Parente (2000) did earlier, one can question the existence of its unique body plan with sterile and fertile laterals occurring in alternation within the same row. Further investigations on this fossil algal genus are requested, requiring the revision of its remaining representatives and its type species.

**Data availability.** No data sets were used in this article.

**Competing interests.** The author declares that he has no conflict of interest.

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